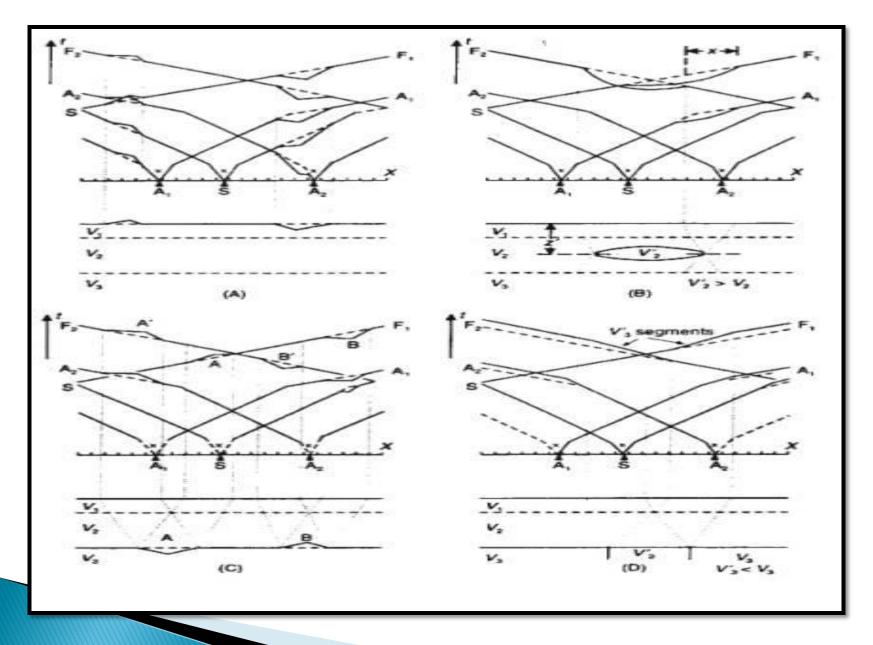
Interpretation of Realistic Travel Time Data

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Interpretation of Realistic Travel Time Data

- With field data it is necessary to examine travel time curves carefully to decide on best method to use:
- How many refraction branches are there, i.e. how many layers?
- II. Are anomalous times due to mispicking or real?
- III. Small anomalies can be ignored, but larger ones require other methods, e.g. Plus-Minus.
- IV. Multiple source positions allow, some inference of depth of anomaly: near-surface anomalies align
- Surface Topography Intervening Velocity Anomaly



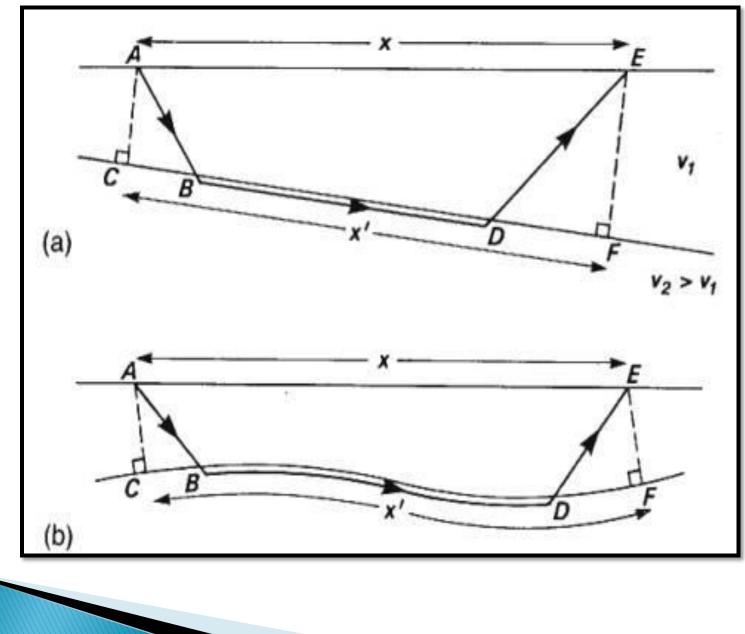
Delay Times

For irregular travel time curves, e.g. due to bedrock topography or glacial fill, much analysis is based on delay times.

Total Delay Time

Difference in travel time along actual ray path and projection of raypath along refracting interface:

$$d = T_{AB} - T_{CF}$$



$$T_{CF} = \frac{CF}{V_2}$$

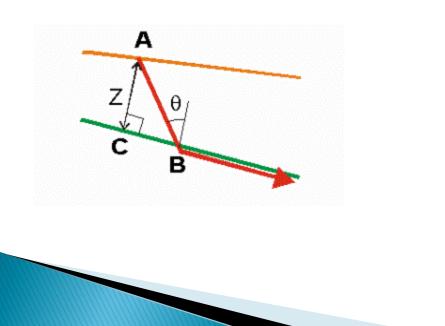
$$\hat{a} = \left(\frac{AB}{V_1} - \frac{CB}{V_2}\right) + \left(\frac{DE}{V_1} - \frac{DF}{V_2}\right) = \hat{a}_S + \hat{a}_G \approx T_{AB} - \frac{x}{V_2}$$

*Total delay time is delay time at shot plus delay time at geophone:

For small dips, can assume x=x'

Refractor Depth from Delay Time

If velocities of both layers are known, then refractor depth at point A can be calculated from delay time at point A:



Solution Using right-angled triangle to get lengths in terms of z:

$$\mathcal{A}_{A} = \frac{Z}{V_{1}\cos\vartheta} - \frac{Z\tan\vartheta}{V_{2}}$$
$$= \frac{Z}{V_{1}\cos\vartheta} \left(1 - \frac{V_{1}\sin\vartheta}{V_{2}}\right)$$

Solution Using Shell's law to express angles in terms of velocities:

$$\mathcal{A}_{A} = \frac{Z}{V_{1} \left(1 - \frac{V_{1}^{2}}{V_{2}^{2}}\right)^{\frac{1}{2}}} \left(1 - \frac{V_{1}^{2}}{V_{2}^{2}}\right)$$

Simplifying:

$$\mathcal{Z}_{A} = \frac{Z(V_{2}^{2} - V_{1}^{2})^{\frac{1}{2}}}{V_{1}V_{2}}$$

Refractor Topography and Refractor Velocity Variation So refractor depth at A is:

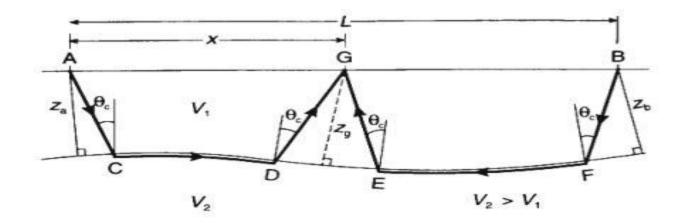
$$Z = \frac{\hat{a}_{A}V_{1}V_{2}}{\left(V_{2}^{2} - V_{1}^{2}\right)^{\frac{1}{2}}}$$

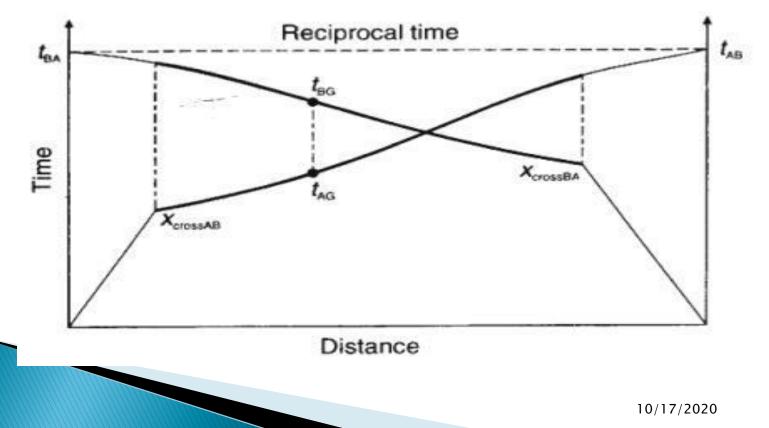
- Additional Hagedoorn's Plus-Minus method used for more complex cases:
- Undulating interfaces
- Changes in refractor velocity along the profile
- Plus–Minus:
- Requires forward and reverse travel times at geophone location to find delay time and refractor velocity at geophone

- Assumes interface is planar between D and E, can result in smoothing of actual topography.
- Assumes dips less than ~10°
- Delay time at G given by:

$$0.5(T_{ACDG} + T_{GEFB} - T_{ACFB})$$

which can be found from observed data.





Plus and Minus Terms

Using previous figure we can write down forward/ reverse travel times:

$$t_{AG} = \frac{x}{V_2(x)} + d\hat{d}_G + d\hat{d}_A$$
$$t_{BG} = \frac{L - x}{V_2(x)} + d\hat{d}_G + d\hat{d}_B$$

Minus Term

Used to determine laterally varying refractor velocity, i.e. $V_2(x)$:

$$T^{-} = t_{AG} - t_{BG} = \frac{2x}{V_2(x)} - \frac{L}{V_2(x)} + \mathcal{A}_A - \mathcal{A}_B$$

•Velocity given by local slope of plot of (T^-) vs. x, distance along profile. Note factor of 2 compared with the plane layer method.

•Velocity may change along profile , so written as V_2 (x) . Different values of V_2 can be used for calculation of interface depth using plus term.

Plus Term

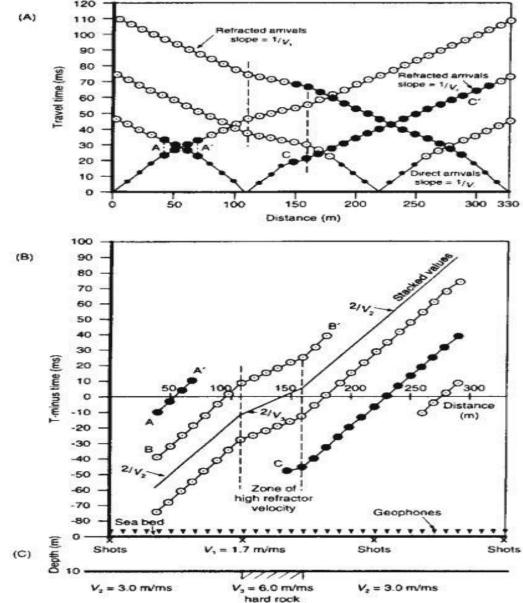
Determines refractor depth at a location from delay time there:

$$T^+ = t_{AG} + t_{BG} - t_{AB} = 2 \mathcal{X}_G$$

So from delay time formula for depth, depth at G given by:

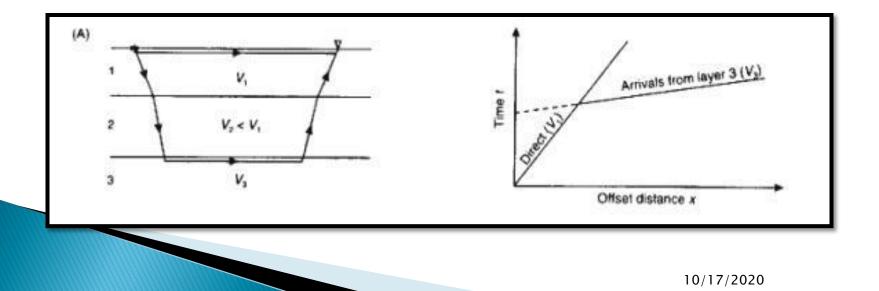
$$z_{G} = \frac{(T^{+})V_{1}V_{2}(G)}{2(V_{2}^{2}(G) - V_{1}^{2})^{\frac{1}{2}}}$$

 Depth can be determined at each geophone location where forward and reverse travel times recorded using V₂ estimated for that position

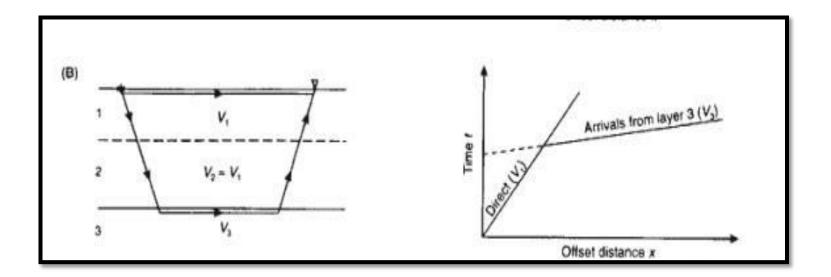


A. Composite travel time distance plots for four different shotsB. Plot of Minus Terms: note lateral changes in refractor velocity

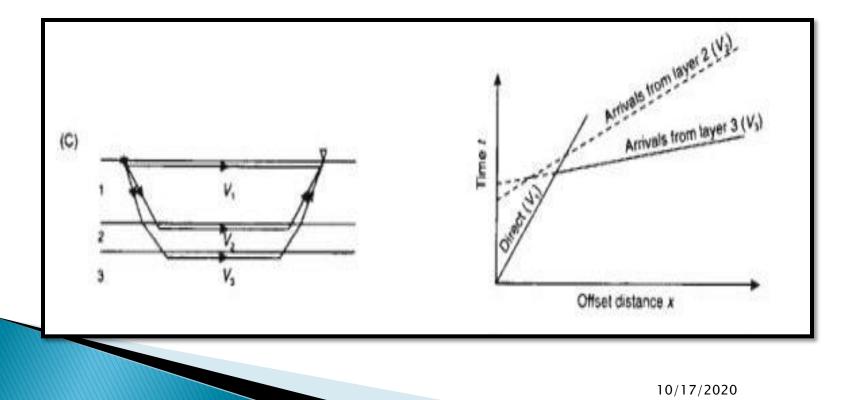
- Layers may not be detected by first arrival analysis:
- A. Velocity inversion produces no critical refraction from layer 2



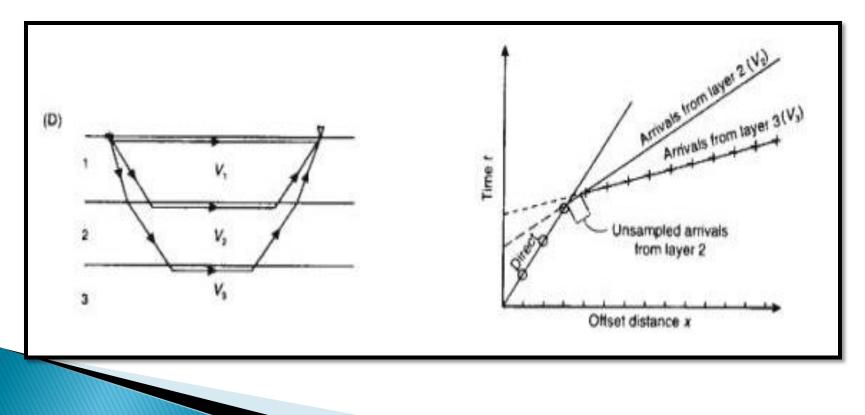
B. Insufficient velocity contrast makes refraction difficult to identify



C. Refraction from thin layer does not become first arrival



D. Geophone spacing too large to identify second refraction



Textbook

Kearey, P., Brooks, M. and Hill, I. (2002) An introduction to Geophysical Exploration. 3rd edition, Blackwell Science Ltd, UK, 261p.